

Global Cloud Climatologies from satellite-based InfraRed Sounders (TOVS, AIRS & IASI)

Claudia Stubenrauch

S. Cros*, A. Guignard, N. Lamquin*,

* until 2010

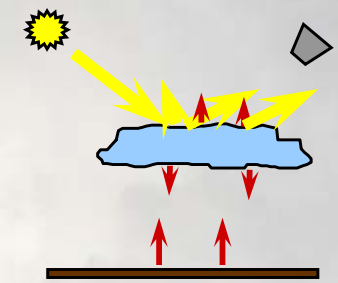
R. Armante, A. Chédin, C. Crevoisier, N. A. Scott,

Atmospheric Radiation Analysis team, ABC(t)

IPSL - Laboratoire de Météorologie Dynamique, France



Cloud properties from space:



1) multi-spectral cloud detection 2) cloud property retrieval

Passive remote sensing (>1980)

info on highest cloud layer
good spatial coverage

- CA (*tot, high, midlevel, low*)
- p/z, T, τ_{VIS} / ϵ_{IR}
- horizontal extension
- bulk microphysical properties

Active (A-Train, >2002)

info on all cloud layers
sampling every 1000km

- Z, τ_{VIS}
- vertical extension
- cloud layering
- microphys. prop. profiles

IR sounders :

- *good spectral resolution -> esp. reliable Ci properties (day & night)*
- *atmospheric T, H₂O profiles (RH) + clouds + aerosols*

A-Train synergy (AIRS-CALIPSO-CloudSat) :

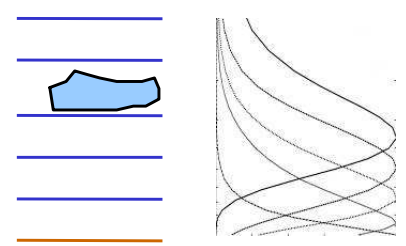
- *choose variables & thresholds for AIRS cloud detection*
- *AIRS cloud height evaluation -> retrieval transfer to IASI*

TOVS-LITE (Stubenrauch et al. JGR 2005)

IR Sounders: TOVS, AIRS, IASI

>1980 NOAA, ≥2002 NASA, ≥2006 CNES

$I_m(\lambda_i)$ along H₂O, CO₂ absorption bands, *good spectral resolution*



Inversion

3I - TOVS (Scott et al. 1999) **AIRS-L2** (Susskind et al. 2003)

- atmospheric temperature & water vapor profiles, T_{surf}

+ atm. transmissivities from TIGR
Thermodynamic Initial Guess Retrieval
 $I_{clr}(\lambda), I_{cld}(p_k, \lambda) \leftarrow$ **4A radiative transfer** \leftarrow radiosondes
 (<http://www.noveltis.fr/4AOP>)

min weighted $\chi_w^2(p_k)$

$$\varepsilon(p_k) = \sum_{i=1}^N \frac{I_m(\lambda_i) - I_{clr}(\lambda_i)}{I_{cld}(p_k, \lambda_i) - I_{clr}(\lambda_i)}$$

no assumption on microphysics !

$\varepsilon_{cld}, p_{cld}$ (Stubenrauch et al. 1996, 1999, 2008, 2010)

{

cirrus emissivities (8 - 12 μ m)

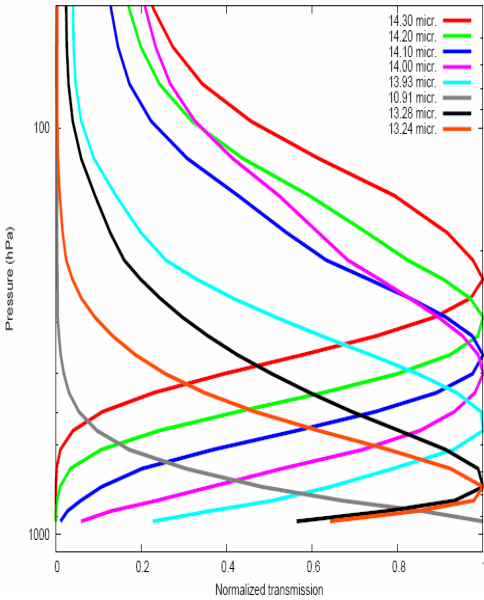
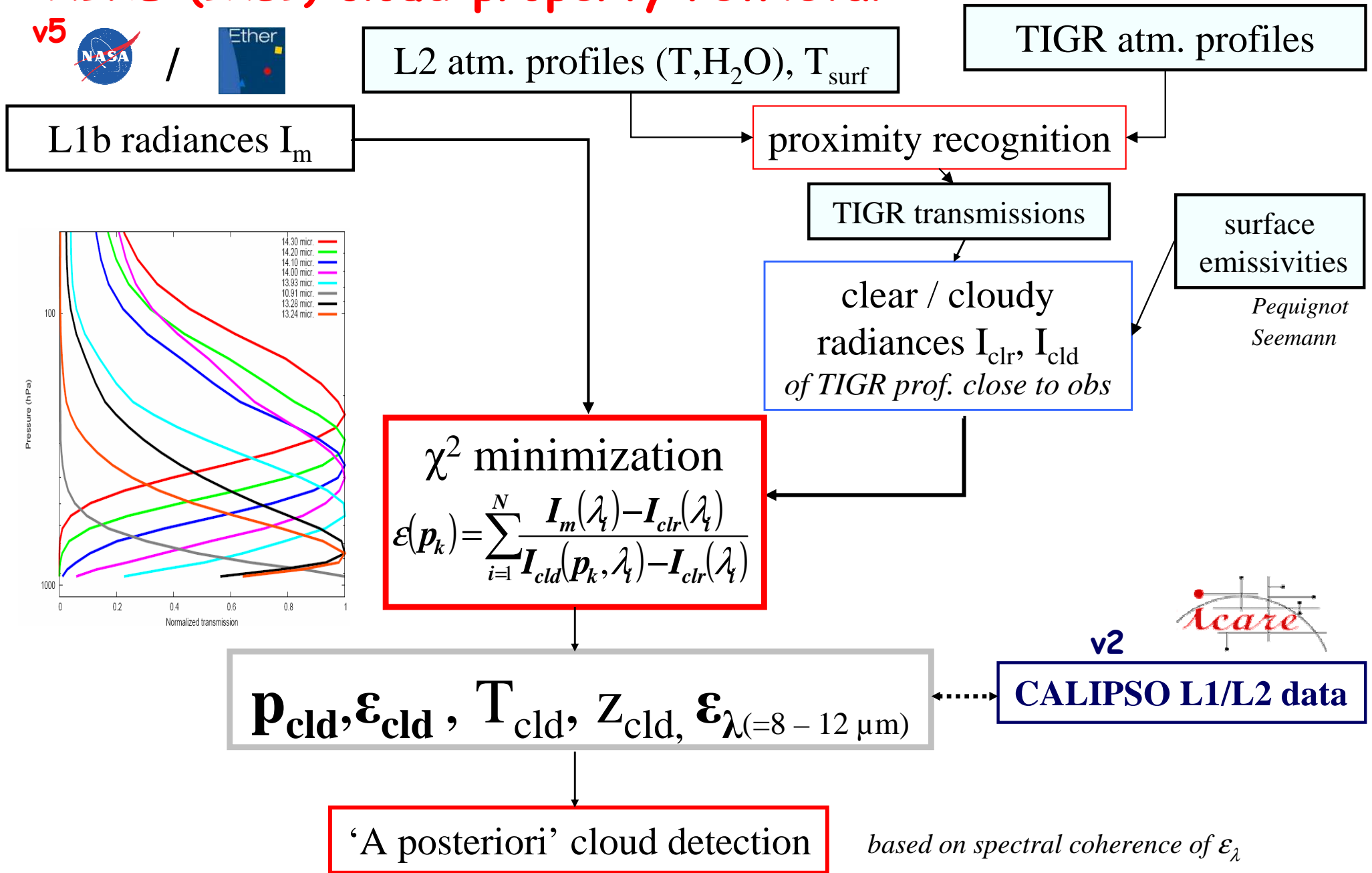
talk Anthony Guignard

simulated LUT

4A-DISORT + SSP of ice crystals

De, IWP (CIRAMOS, Rädcl et al. 2003, Stubenrauch et al. 2004) *Mitchell, Baran*

AIRS (IASI) cloud property retrieval:



cloud property retrieval

TOVS Path-B

AIRS-LMD

Cloud detection: **multi-channel (+MSU)**
(\lt NOAA15: MSU channels θ_v corrected)

$\sigma(\epsilon_{\text{cld}}(\lambda))/\epsilon_{\text{cld}}$ & $T_{\text{cld}} - T_{\text{surf}}$
a posteriori

Spatial resolution:

T profile: 100 km (MSU)

45 km (AMSU)

cloud detection: 17 km

13 km

cloud retrieval: av cldy 100km

13 km

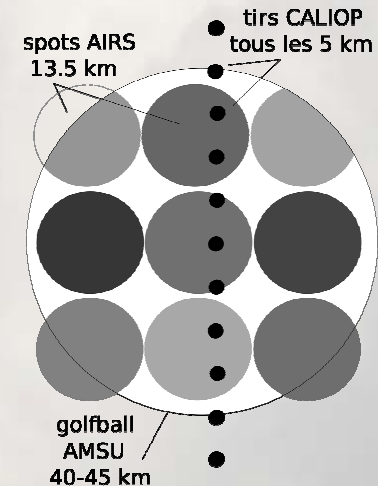
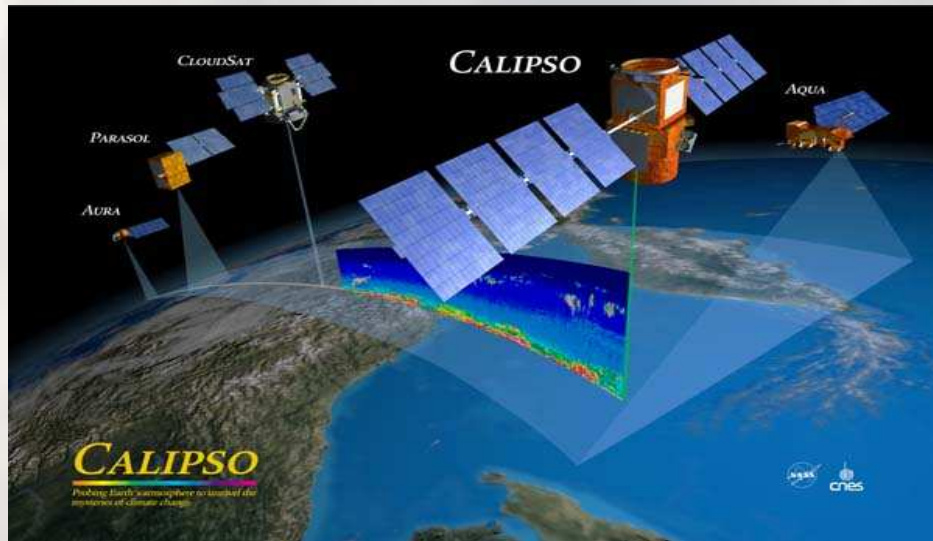
Observation time: 7h30, 19h30 - 1h30, 13h30

1h30, 13h30

Sampling: **every MSU footprint**

every AMSU footprint

A-Train Synergy: AIRS-CALIPSO-CloudSat

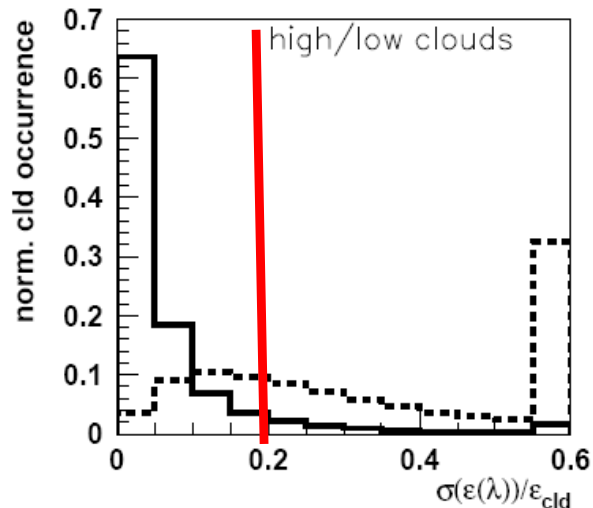


- 1) Development of cloud detection method
- 2) Evaluation of AIRS cloud height
- 3) Vertical extent (Δz) of high opaque clouds / Ci / thin Ci
- 4) Vertical insight into high opaque clouds / Ci / thin Ci
- 5) Cloud height relative to tropopause
- 6) Multiple scattering correction of Ci lidar signal



1) Development of ‘a posteriori’ cloud detection

CALIPSO sampling: (5 km x 0.07 km) in (13.5 km x 13.5 km)



CALIPSO samples indicate reality

— CALIPSO cloudy
 - - - CALIPSO clear sky

cloudy:

$$\epsilon_{cld} > 0.05$$

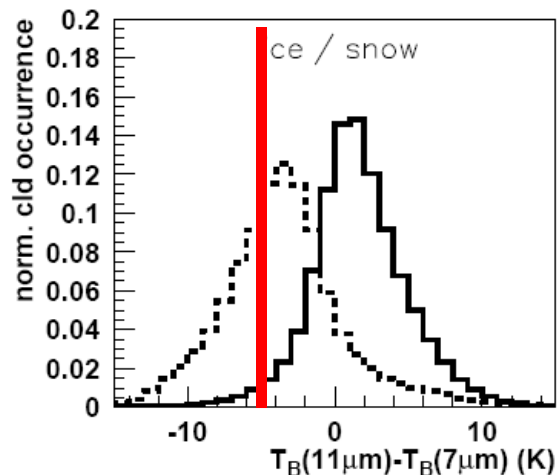
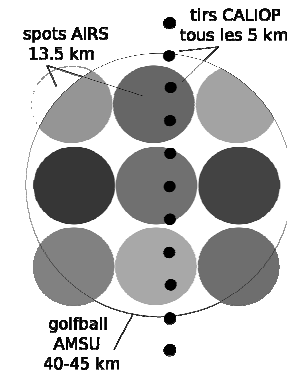
$$\sigma(\epsilon_{\lambda}) / \epsilon_{cld} < 0.2 \quad (9 \mu m < \lambda < 12 \mu m)$$

$$\& \text{ over land: } T_{cld} - T_{surf} > -3K$$

ice/snow (L2, detected by microwave):

$$\sigma(\epsilon_{\lambda}) / \epsilon_{cld} < 0.3 \ \& \ T_B(11\mu m) - T_B(7\mu m) > -5K$$

similar to MODIS-ST (Frey et al. 2008)



Cloud detection comparison & sensitivity:

agreement with CALIPSO: 85 (80) % over ocean (sea ice)
80 (73) % over land (snow)

what is a cloudy AIRS footprint?

global CA : 66 % (*0% for not cloudy*) - 74% (*30% for not cloudy*)
HCA/CA: 40% - 36% **LCA/CA:** 42% - 47%

agreement with LMD multi-spectral cloud detection (ΔT_B):

(for determination of trace gases, aerosols & surface emissivity)

70% over ocean (30°N-30°S)

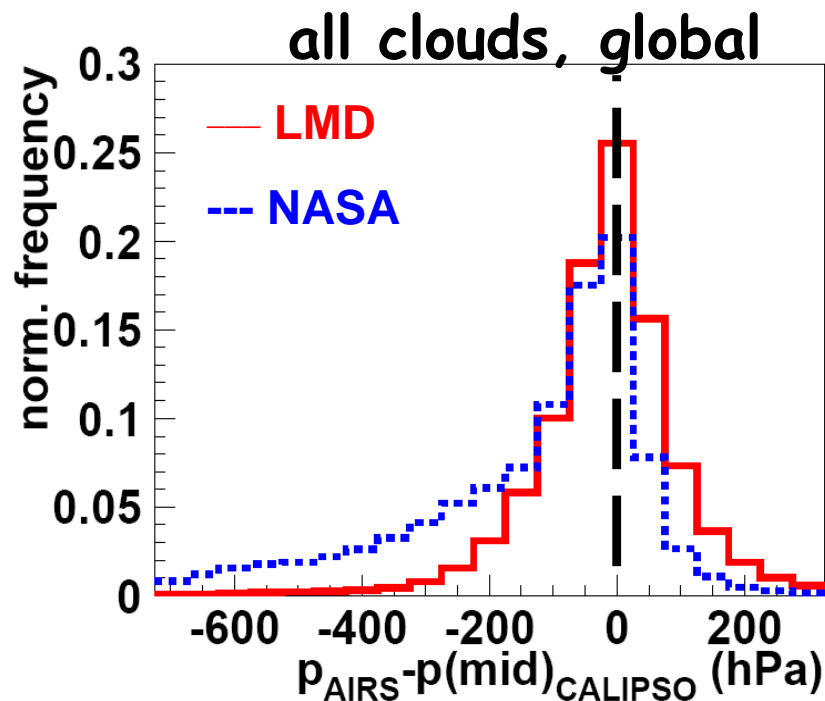
72% over land (30°N-30°S)

Multi-spectral cloud detection about 15% more cloudy footprints
(lowlevel cloud fields)



2) Evaluation of AIRS cloud height with CALIPSO

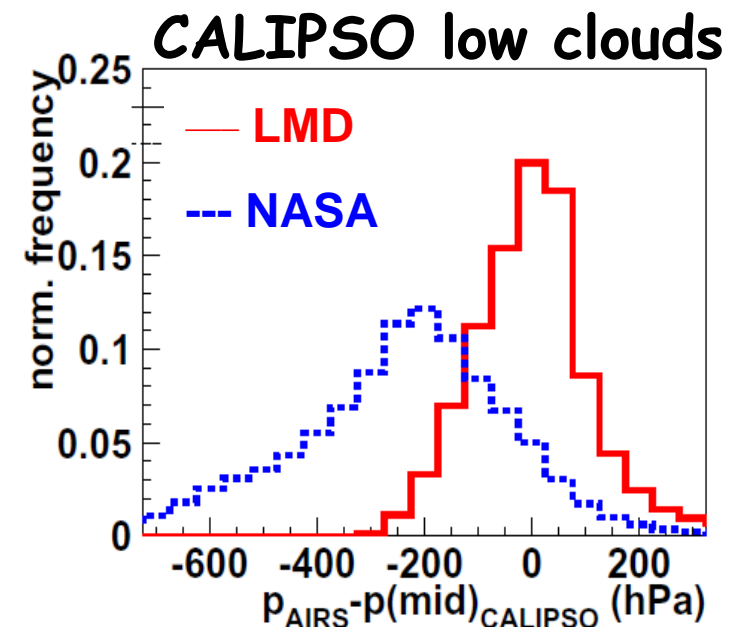
(highest cloud, detected at $\leq 5\text{km}$)



Stubenrauch et al. JGR 2008, ACPD 2010

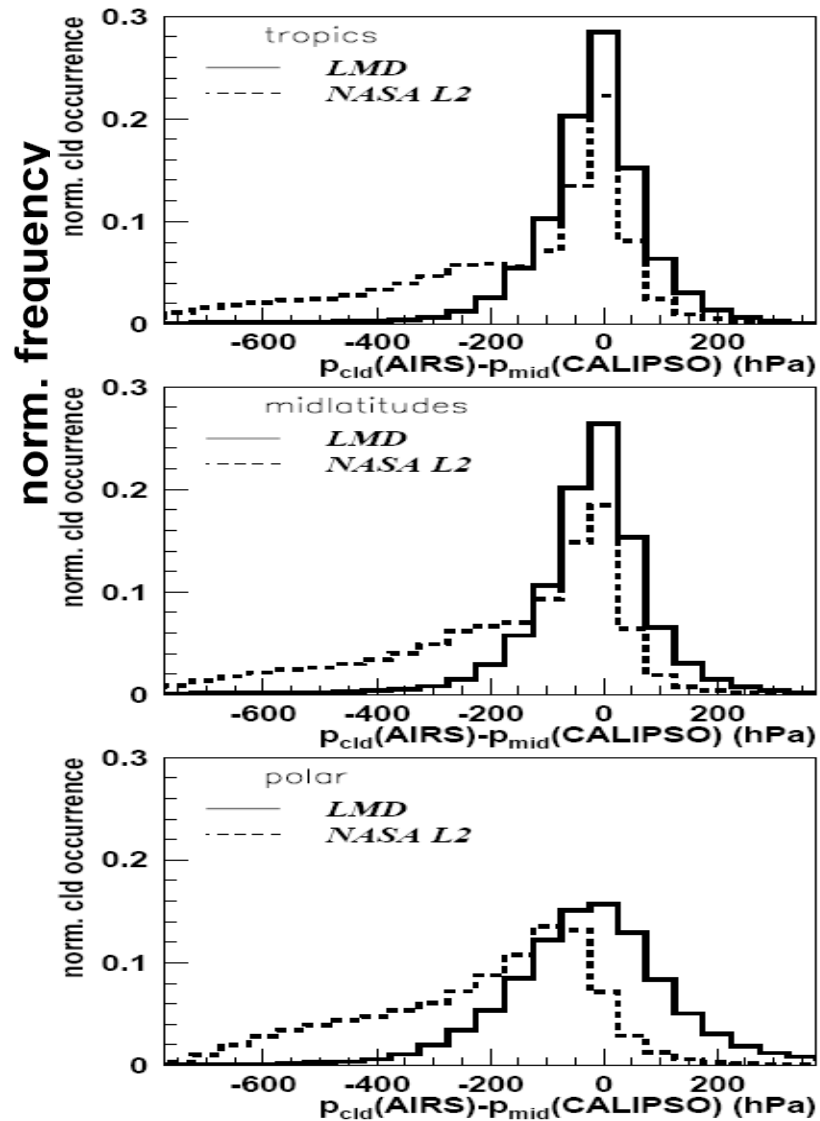
in agreement with Kahn et al. 2008

good agreement with CALIPSO cld midlevel
(or pos. of max. backscatter)
properties also depend on retrieval method



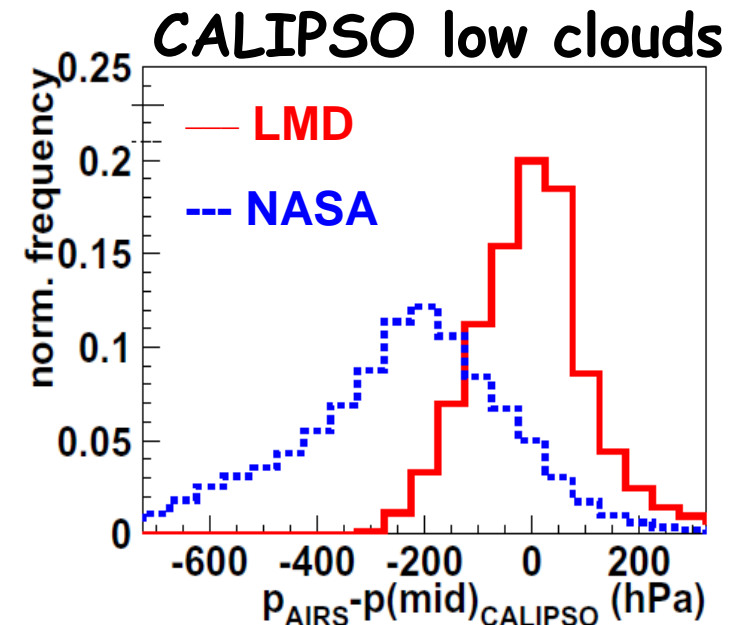


2) Evaluation of AIRS cloud height with CALIPSO



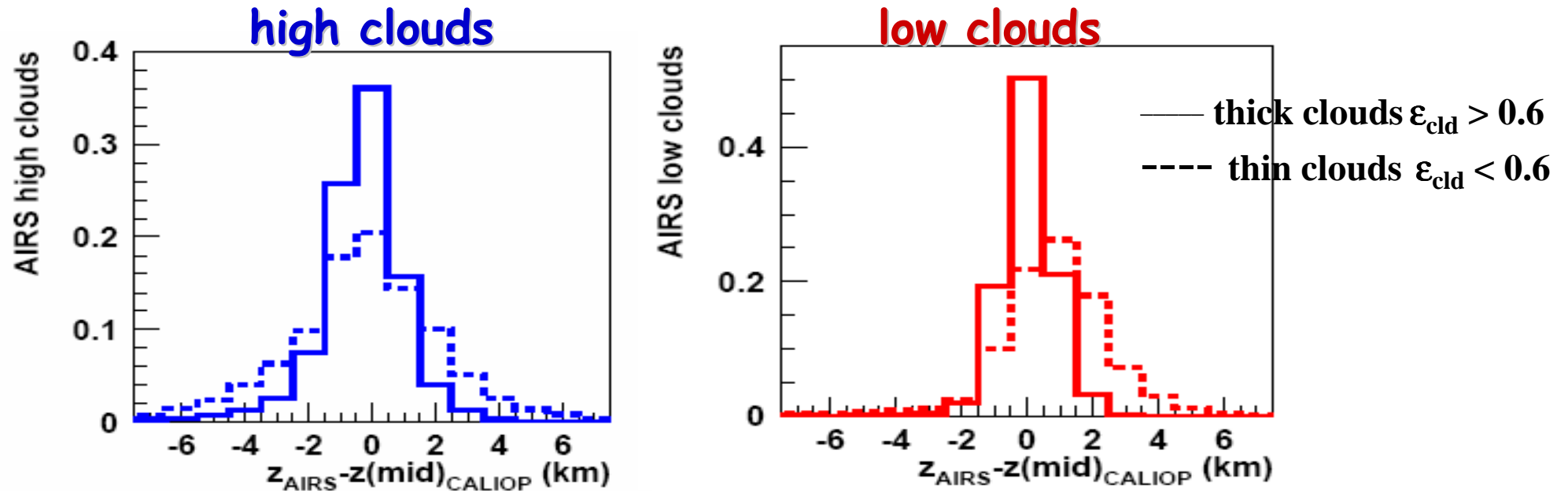
(highest cloud, detected at $\leq 5km$)

good agreement with CALIPSO cld midlevel
(or pos. of max. backscatter)
properties also depend on retrieval method



2) Evaluation of AIRS cloud height with CALIPSO

Stubenrauch et al., JGR 2008



good agreement with CALIPSO midlevel of cloud (*highest with $\tau > 0.1$*)
slightly broader distributions for optically thinner clouds, but no bias
sampling: (5 km x 0.07 km) in (13.5 km x 13.5 km)

$\Delta z_{\text{mid}}(\text{AIRS-CALIPSO}) \pm 1.5 \text{ km:}$

High: 51% 55% 66%

Low: 70% 74% 80%

hghst / hghst w $\tau > 0.1$ / closest layer

$\Delta p_{\text{mid}}(\text{AIRS-CALIPSO}) \pm 75 \text{ hPa:}$

High: 72% 81% (thick); 63% (thin)

Low: 59% 69% ; 38%



A-Train: synergy of passive & active instruments

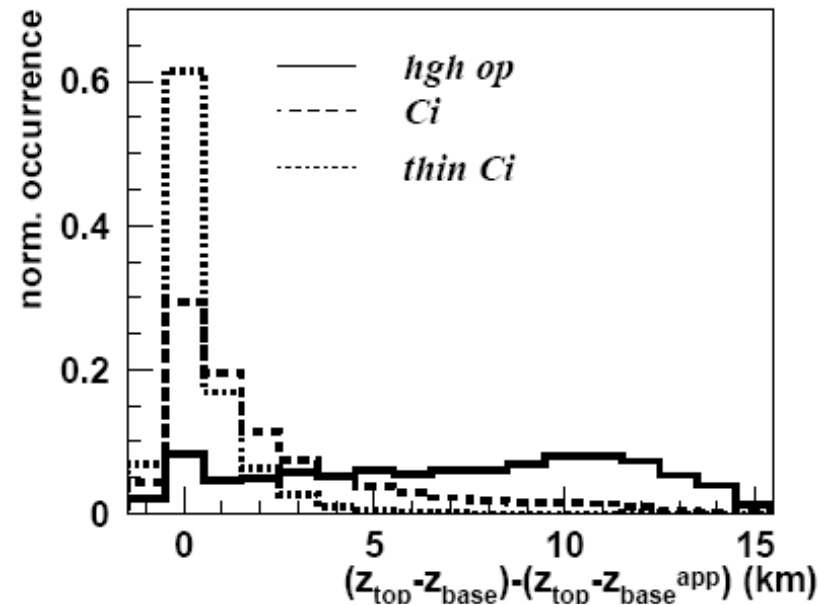
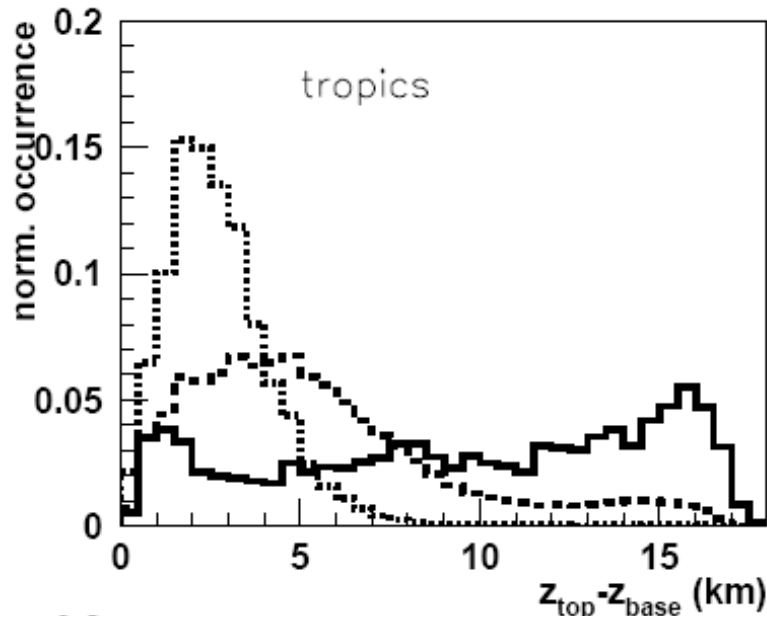
3) Vertical extent (Δz) of high opaque clouds / Ci / thin Ci

AIRS: cloud type

CALIPSO: ‘apparent’ geometrical cloud thickness

CloudSat: real geometrical cloud thickness

} *Winker / Mace et al.2009*
GEOPROF data
Cloudsat.cira.colostate.edu



- $\Delta z(\text{thin Ci}) < \Delta z(\text{Ci}) < \Delta z(\text{hgh op})$
- real Δz much larger than apparent Δz for high opaque cloud
- good quality of AIRS cloud type identification

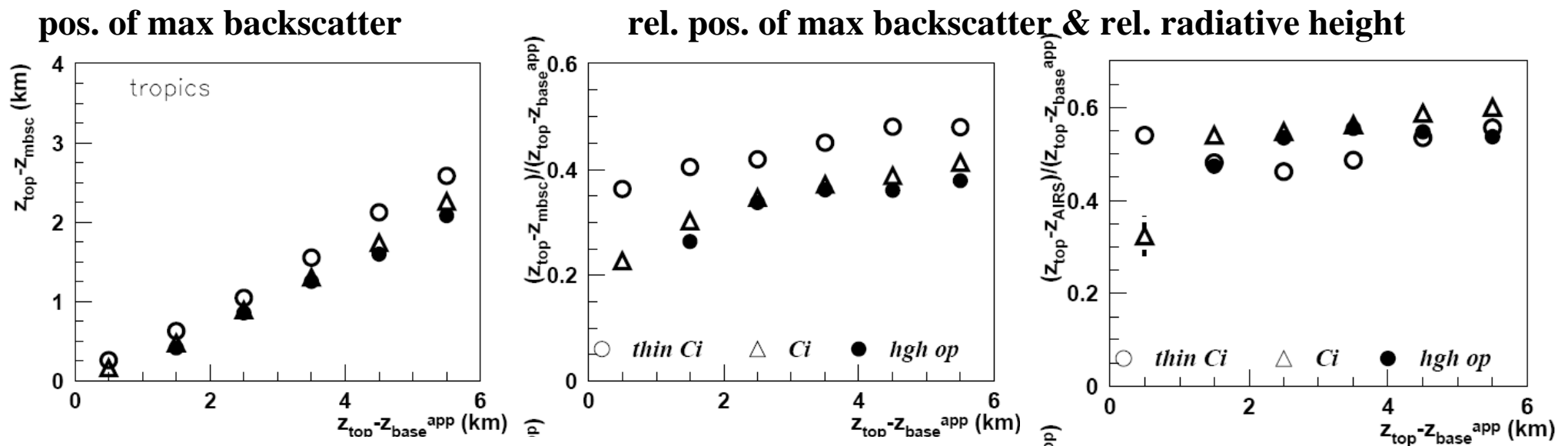


4) Vertical insight into high opaque clouds / Ci / thin Ci

AIRS: cloud type

CALIPSO: 'apparent' geometrical cloud thickness, position of max. backscatter

CloudSat: real geometrical cloud thickness

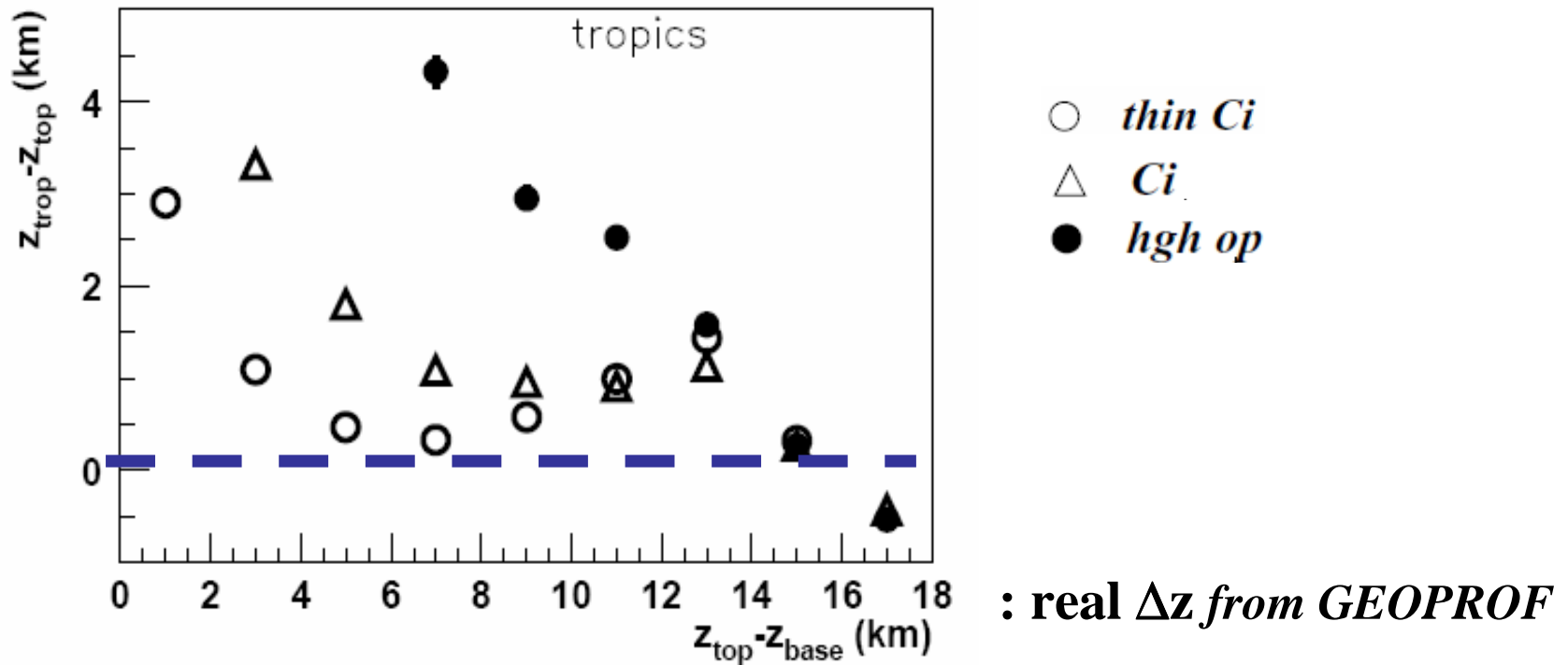


- position of max backscatter depends on 'apparent' Δz & can reach 2 km below cloud top, even for high opaque clouds
- rel. position less dependent, 1/3 - 1/2 below top (thin Ci)
- 'radiative' height lies about 1/2 below top, for all cloud types



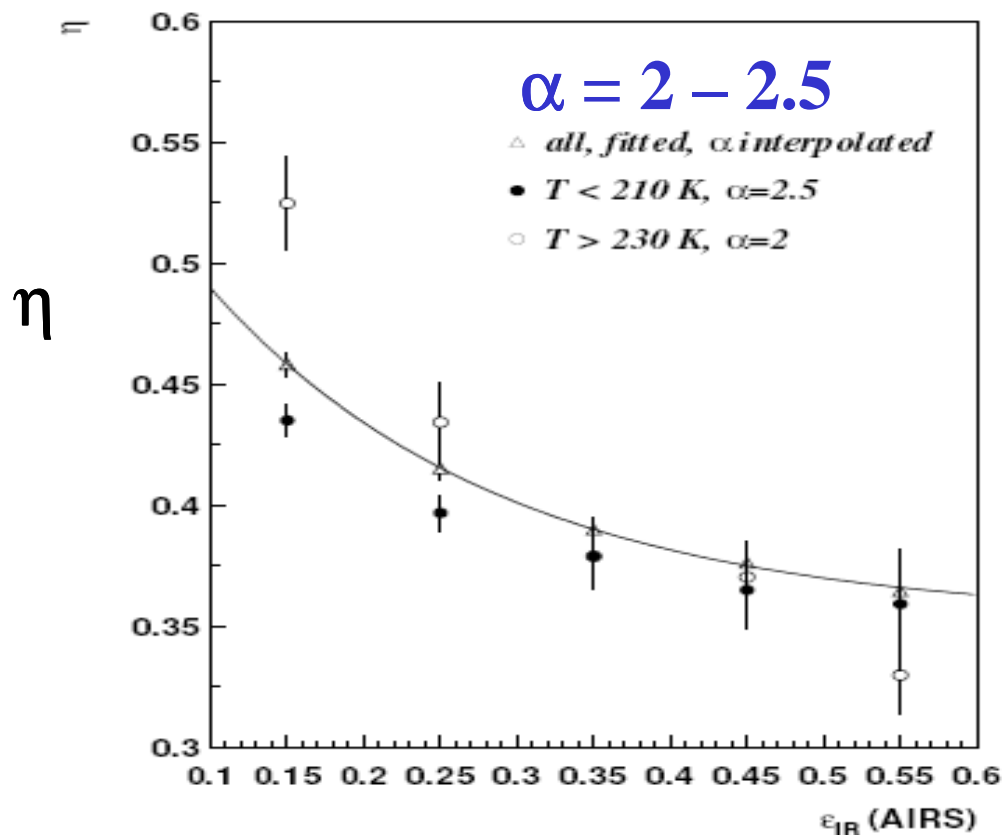
A-Train: synergy of passive & active instruments

5) Cloud height relative to tropopause



Tropics: only the very thickest opaque clouds (& surrounding anvils) penetrate stratosphere
Rossow & Pearl 2007: larger, organized, convective systems penetrate

6) Multiple scattering correction of cirrus lidar signal



Lamquin et al. JGR, 2008

$$\epsilon_{\text{IR}} = 1 - \exp(-\tau_{\text{VIS}}/\alpha), \quad \alpha \cong 2 - 2.5$$

α increasing with decreasing crystal size

(Sassen & Comstock 2001)

D_e increases with T_{cld} , (Stubenrauch et al. 2004)

$\Rightarrow \alpha = f(T_{\text{cld}})$

$$\epsilon_{\text{IR}} = 1 - \exp(-\tau_{\text{VIS}}/\alpha)$$

$$\tau_{\text{VIS}} = \tau_{\text{LIDAR}} / \eta$$

$$\eta = 0.47 - 0.37$$

$$\text{CALIPSO L2: } \eta = 0.6$$

slightly more multiple scattering for opt. thicker cirrus

Relative amount of high clouds *according to:*

CALIPSO (*with-without subvis Ci*) / AIRS - TOVSB / HIRS / ISCCP VIS+IR / ISCCP IR

51%	-	45%	/	39%	-	41%	/	49%	/	34%	/	22%	NH midlatitudes
66%	-	57%	/	58%	-	62%	/	60%	/	46%	/	30%	tropics
40%	-	34%	/	28%	-	30%	/	49%	/	26%	/	17%	SH midlatitudes

Relative amount of low clouds *according to:*

CALIPSO (*with-without subvis Ci*) / AIRS - TOVSB / HIRS / ISCCP VIS+IR / ISCCP IR

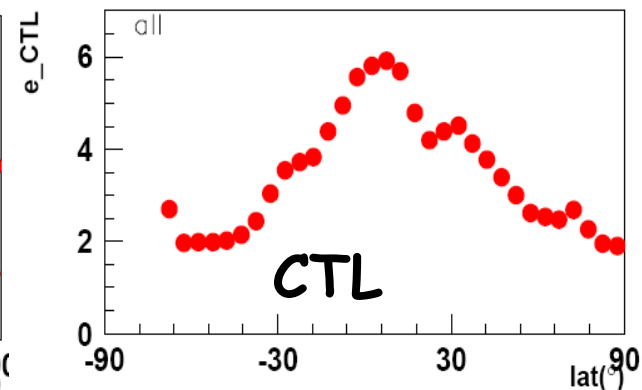
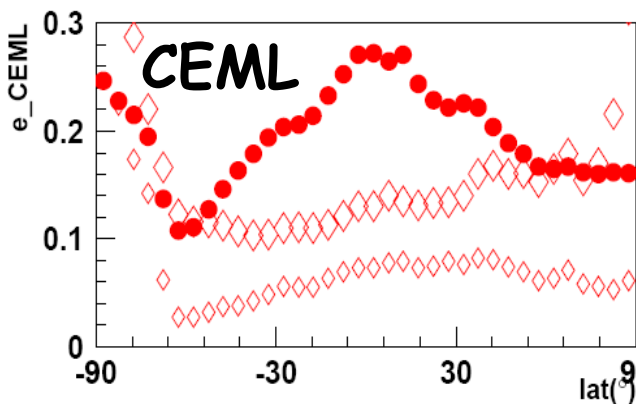
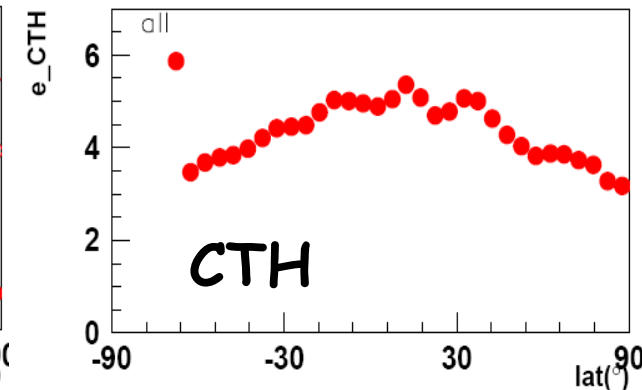
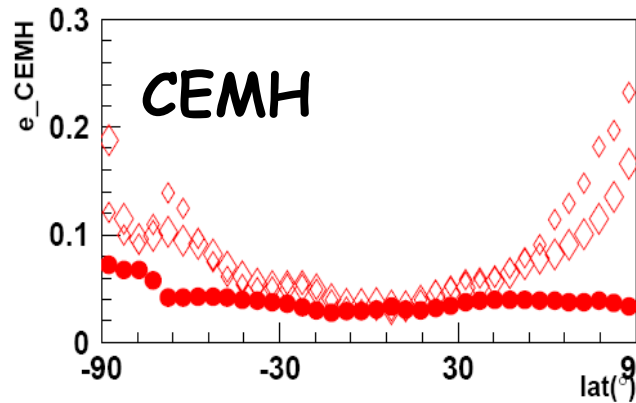
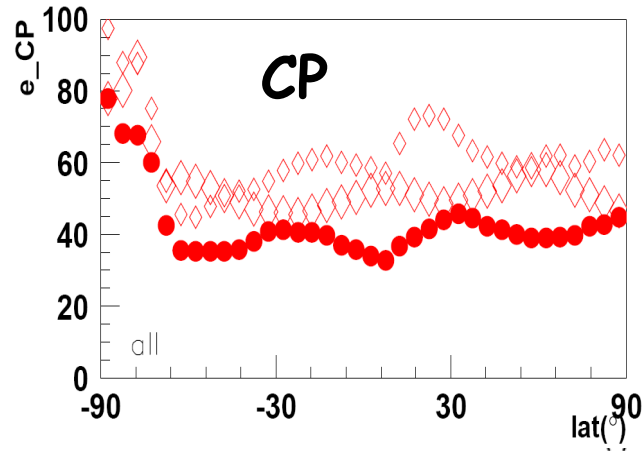
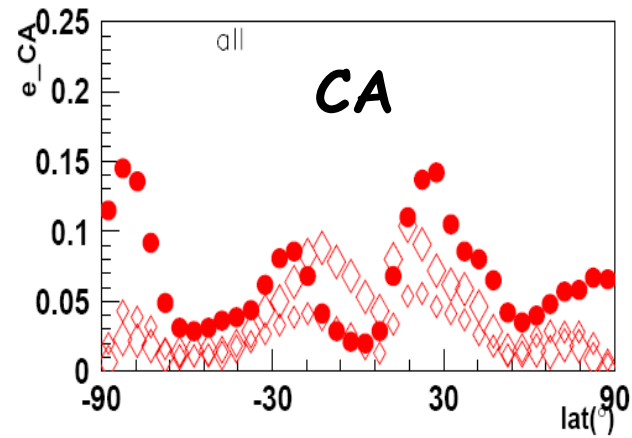
34%	-	36%	/	39%	-	39%	/	33%	/	38%	/	41%	NH midlatitudes
28%	-	34%	/	31%	-	34%	/	30%	/	32%	/	45%	tropics
51%	-	52%	/	51%	-	52%	/	29%	/	49%	/	48%	SH midlatitudes

AIRS-LMD, TOVSB in good agreement with CALIPSO (without subvisible cirrus)

HIRS no difference between NH & SH midlatitudes

ISCCP same latitudinal behaviour, but underestimation of Ci (5-15%)

Uncertainties



June 2010

GEWEX CA, Berlin

Estimation:

CA: $CA(\epsilon_{\text{cld}} < 0.05)$

largest over deserts & Antarctica

CP, CT, CEM:

difference between $\min \chi^2$ & next smallest χ^2

CP: ~40-50 hPa, slightly smaller for AIRS

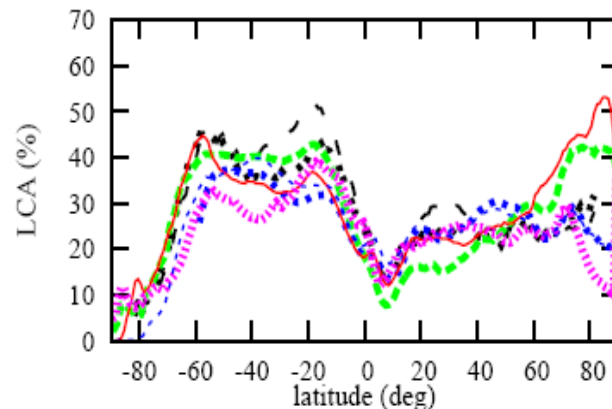
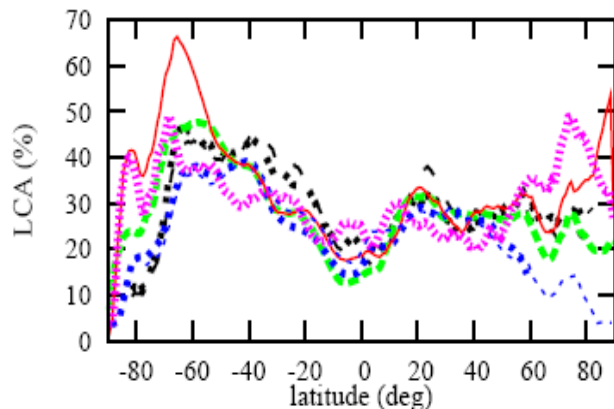
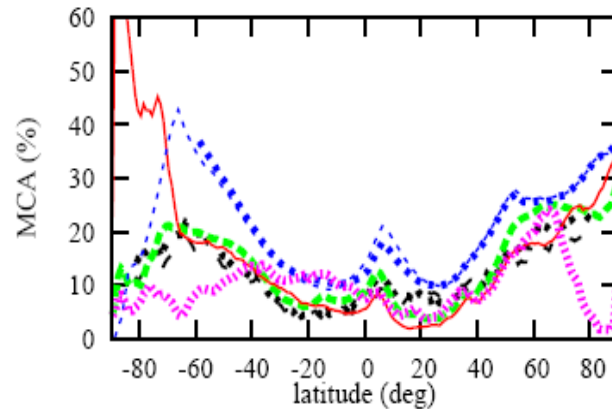
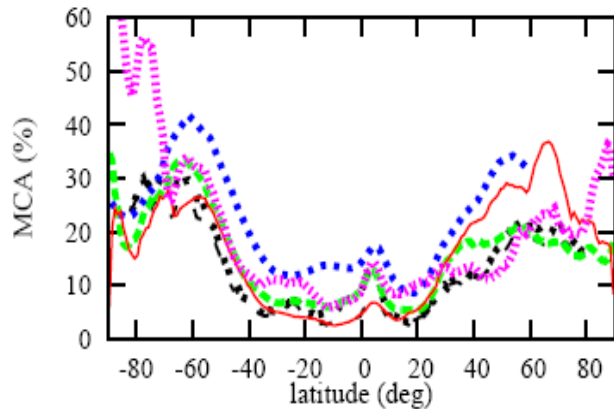
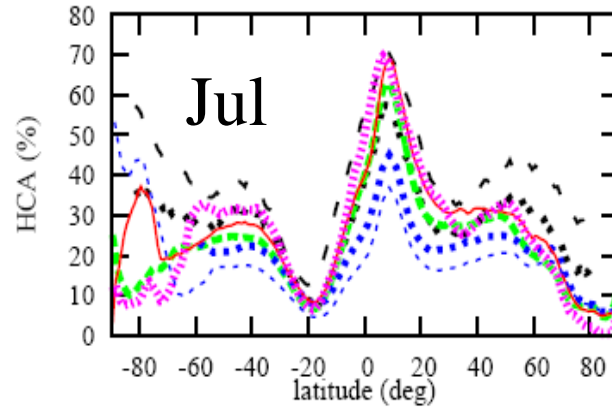
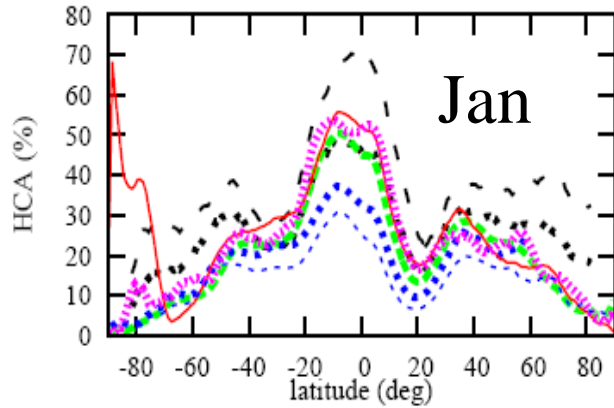
High clouds:

CT: 4K, CEM 0.05

Low clouds:

CT: 2-4K, CEM: 0.1-0.25 slightly smaller for TOVS

Zonal averages of HCA, MCA and LCA



CALIPSO 2007-2008
ISCCP 1984-2004
TOVS-B 1987-1995
AIRS 2003-2008
IASI 2008

*IASI preliminary:
 use of av. AIRS L2 atm. profiles*

latitudinal behavior similar
HCA/CA:
CALIPSO highest
TOVS ~ AIRS ~ IASI
ISCCP lowest
LCA/CA:
agreement,
except at higher latitudes

CALIPSO	- - -	ISCCP
CALIPSO 5km	TOVS-B	————
AIRS-LMD	IASI

TOVS Path-B (1987-1994) & AIRS-LMD (2003-2009) cloud climatologies:

❖ IR sounders passive instruments most sensitive to cirrus

5-15% more Ci than ISCCP VIS+IR

A-Train: unique possibility to evaluate IR sounder retrieval & to give insight into vertical structure of different cloud types :

LMD retrieval:

❖ p_{cld} corresponds to midlevel of apparent cloud depth
(slightly below height of max backscatter)

❖ Uncertainty estimation from χ^2 :

CP : 40 - 50 hPa

High clouds: CT: 4 K, CEM: 0.05

Low clouds: CT: 2-4K, CEM: 0.1-0.25



This work was supported by CNRS and CNES.

We also thank all Science teams as well as the engineers and space agencies for their efforts and cooperation in providing the data.